Study of Zircon Crystals: An Implication to Determine of Source and Temperature of Crystallization in Turkeh Dareh Pluton, NW Iran

M. Jamshidibadr*

Department of Geology, Faculty of Sciences, University of Payam Noor, Tehran, Islamic Republic of Iran

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Abstract
Turkeh Dareh pluton is outcrop in NW of Iran and it is one of the intrusions in Sanandaj-Sirjan zone. Shaped crystals of this pluton distinct oscillatory zoning narrow and close together zoning in zircons indicate parental magma richness of zirconium. Factors affecting the shape of the zircon crystals are the composition, possibly the temperature of the crystallization, water content in magma and origin of the magma, therefore, were studied morphology of zircon crystals of Turkeh Dareh pluton. Morphologically, most of Turkeh Dareh zircons are located in range of S13 and S14 and in lesser extent in the range of S18, S19 and S20. The minimum zircons crystallization temperature was calculated 750 to 840 C° based on the morphology of zircon, zircon saturation and the whole rock geochemistry. The origin of intrusion magma of Turkeh Dareh is I-type and calc-alkaline that these results are consistent with results of field observations, mineralogy and geochemistry of this pluton.

Keywords: Morphology of zircon; Thermometer; Turkeh Dareh Pluton; Sanandaj-Sirjan zone; Iran.

Introduction
The resistance of zircon to chemical and mechanical processes as well as its high melting point is so high that it can remain intact and unchanged within the earth crust for millions of years, and therefore it can survive from processes such as weathering, transportation, high temperature metamorphism and even partial melting. Morphology of zircons is closely related to the origin, development, and chemical composition and geological position of its parental magma [1-14]. Zircon crystallizes in the tetrahedral system in the form of two-way prismatic crystals with length to width ratio of one to five. The high ratio of length to width in this mineral is signs of rapid crystallization of parental magma [15].

In addition, zircon has great tendency to trace elements, which are distributed in different zones representing changes in elements during magmatic evolution [16]. Pupin [2] ranked zircon crystals based on relative growth of prismatic shapes {100} to {110} and pyramidal {211} to {101}. Pupin [2] suggested that the zircon typology is a function of chemistry and crystallization temperature of a magma from which zircon crystallizes. He further shown that the relative growth of pyramidal shapes with a chemical composition (ratio Al / (Na + K) or indexes A and

* Corresponding author: Tel: +982645383686; Fax: +982645383244; Email: m_jamshidi@pnu.ac.ir
relative growth of pyramidal shapes are directly related to the crystallization temperature. In other words, zircons crystallized from aluminous liquids have pyramids \{211\} and zircons crystallized in the \{101\} in alkaline conditions have pyramid with developed levels. According to Pupin [2], magma temperature is the most important factor controlling relative growth of various prismatic shapes in zircon crystals. Therefore, crystallized shape of zircon can be used as thermometer.

However, as Vavra [17] interpreted growth rate and size of zircon crystals are controlled oversaturation of melt from ZrSiO$_4$ and the amount of trace elements and these factors are more important than temperature. According to him, pyramids morphology (eg \{211\} versus \{101\} is basically controlled external trace elements, while the prism form is determined the supersaturation degree of ZrSiO$_4$. According Benisek & Finger [16], the amount of uranium and equation 1 are important factor of growth in prisms \{110\} of zircon. That is why the granitic magmas containing low levels of U, Th, Y, REE and P (compared to Zr) crystallize zircons with large prisms \{100\}.

\[(Y, \text{REE})^{3+} + P^{5+} = Zr^{4+} + Si^{4+}\] (Equation 1)

Investigation morphology of zircon crystals in granite rocks, primary magma source was introduced different researchers for e.g [2, 18, 19]. They distinguished magma sources of three groups of granites-I (tholeitic and alkaline granite high temperature), granites-S (moderate temperature granites derived from shells or aluminus) and two-vein granites (low temperature granites of calcic-alkaline and sub-alkaline series.

In this study, zircon was used to investigate the origin of the magma and parental magma thermometer of Turkhe Dareh pluton, due to its specific characteristics such as chemical and mechanical characteristics and high melting point. Turkhe Dareh intrusion has outcrop in the North West of Iran and among cities of Shahindej and Takab and it is among the intrusions of Sanandaj-Sirjan zone introduced in the geological map of 1:100,000 in Shahin Dej [20] and Chapan [21] (Fig. 1). The age of Turkhe Dareh intrusion is 59.0± 2.7 based on U-Pb dating of zircon reported by [22] and the age of regional metamorphism (schist) were dated from a garnet-schist using the yielded a $^{238}$U/$^{206}$Pb age of monazites 61±8 ma that reported by [23]. Using morphology of zircon separated from intrusion rocks of Turkhe Dareh and field features,
geochemistry and geomorphology of this pluton was discussed so that the relationship between these studies can be a model for future studies. The study of zircon morphology used only for igneous zircons, its not be used for metamorphic zircons [2, 18, 19].

Materials and Methods

The zircon separation was carried out in the laboratories of the University of Adelaide in Australia. Zircons were separated using conventional methods that include crushing, sieving, magnetic separation and floatation. More than fifty zircon grains were handpicked under a binocular microscope. The zircons were then set in synthetic resin mounts, polished and cleaned in a warm HNO₃ ultrasonic bath. Cathodoluminescence (CL) and back-scattered electron (BSE) imaging were carried out to help characterize any compositional variation within individual zircons (Figs. 2 & 3). After microscopic studies, 8 samples were selected for whole rock geochemistry studies. Analysis was conducted using XRF method in Tehran Tarbiat Moallem (teacher training) University (Table 1).

General geology

Turkeh Dareh pluton has light to dark gray color (Figs. 4a & b) and in some parts centralization of minerals in one part and dark minerals in other part caused layer in this mass leading to light and dark gray in this part (Fig. 4c). Due to centralization of dark minerals in some parts of different samples with different colors to differentiate the zircon and petrographic and geochemical studies, sample was taken and examined. In microscopic studies, this pluton had granular texture and in terms of mineralogy it has euhedral to subhedral plagioclase with polysynthetic

![Figure 2](image.png)

Figure 2. A) The location of various types of zircon crystals within petrogenetic categorization of Pupin [2], B), C) and D) Zircon crystals of Turkeh Dareh pluton and their classification based on morphology after [2].
twinning, anhedral quartz, and alkali feldspar and mafic minerals in the pluton are biotite and amphibole. The accessory minerals of this pluton include titanite that are perfectly shaped and dark brown color, zircon and apatite (Figs. 5a, b & c). From mineralogy perspective, dark and light parts of this pluton are not different and only modal percentage of felsic and mafic minerals in this part is different. Light parts have high frequency of felsic minerals than mafic minerals. In naming Streckeisen & LeMaitre [24], modal of this pluton can be seen in the range of granodiorite and monzodiorite (Fig. 6a).

Table 1. XRF analyses on Turkeh Dareh granitoids [21] and calculation of M, degree of saturation and saturation temperature of zircon

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<td>70.73</td>
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<td>0.4</td>
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<td>0.69</td>
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<td>787.70</td>
<td>782.32</td>
<td>702.53</td>
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Figure 3. A), B), C) and D) Cathodoluminescence images of zircon grains showing long, euhedral and oscillatory zoning (the scale bar is 200 μm)[21].
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Figure 4. Turkeh Dareh pluton outcrop. A) Light gray color including quartz, plagioclase, K-feldspar, and some amphibole and biotite, B) magmatic banding from Turkeh Dareh pluton, C) centralization of hornblende and biotite in some parts of Turkeh Dareh pluton that shows dark gray color.

Figure 5. A) microphotograph from a light gray area of Turkeh Dare pluton granitoids. Felsic minerals are more abundant than (B) microphotograph. myrmekitic texture seen in centre of photo, C) microphotograph of dark band from Turkeh Dare pluton granitoids. The rock contains quartz, plagioclase, K-feldspar, hornblende, biotite, Titanite. Note that the hornblende and biotite are oriented.
Results

Whole rock geochemistry

According to field and petrographic studies, the Turkeh Dareh pluton has dark and light parts. In the light parts, the frequency of Na₂O, K₂O, and SiO₂ is greater than in the dark parts due to the centralization of felsic minerals. Frequency percentage of Fe₂O₃, MgO, and CaO in samples related to dark parts is higher than in the light parts (table 1). In the multi-cationic R₁-R₂ diagram [25], samples of the Turkeh Dareh are mainly in the range of granodiorite and tonalite (Fig. 6b). Regarding saturation of aluminum [26], samples of the Turkeh Dareh intrusions have values of A/CNK < 1.1, and they are in the range of I-type granites and magmatic series of the Turkeh Dareh calc-alkaline with high and moderate potassium (Figs. 6c & d). The Turkeh Dareh pluton has weak anomaly (Eu/Eu* = 0.95-1.48) related to the absence or slight subtraction of plagioclase. The negative anomalies in Nb and Ta in normalized graphs indicate that their formation is in subduction zones that are the result of fluids and melts due to subducted lithospheric with metasomatism of mantle wedge above themselves [28]. In the tectonic environment separation diagrams, the Turkeh Dareh pluton samples are in the range of volcanic arc [22].
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Morphology of zircon crystals
Zircon crystals separation of various samples of Turkeh Dareh intrusion that had different frequency in mafic and felsic minerals was conducted. Separated zircons are honey yellow to colorless and transparent to translucent with zoning. In terms of morphology, most of zircon crystals in the classification of Pupin [2] are in the range of S14 and S13 and to a lesser extent in the range of S18, S19 and S20 (Figs. 2, 3 & 7). With regard to the distribution of zircons in the diagram provided from Pupin [2], the mean values of Alkalinity Index (IA) and Temperature Index (IT) that were calculated distribution of studied zircons in the diagram of Pupin [2] are respectively 461.8 and 445.5 (Tables 2a & b). In addition, to calculate typological evolutionary trend, a line with slope of ST/SA (standard of deviation of index T and standard deviation of index A) has been drawn from intersection of IA and IT that is equals to tangent of angle of two axes T.E.T and I.A [2]. The typological evolutionary trend measured for studied zircons has angle of 46.02 degrees with axis I.A. Therefore, values IA, IT zircon typological trend of studied granitoids are similar to zircons with calc-alkaline magmatic origin and they are at the range of granodiorite, monzonite, and monzogranites (Fig. 7).

Discussion

Geo-thermometry of crystals of zircon
Based on crystals of zircon [2] and equations provided from [6, 7] and using geochemical data, we can estimate the minimum temperature of zircon formation in a magmatic source. Using the above-mentioned cases, minimum temperature of Turkeh Dareh intrusion rocks was calculated.

Geo-thermometry of Turkeh Dareh pluton based on crystal shapes of zircon
Based on the morphology of zircon crystal, zircon crystallization temperature of Turkeh Dareh pluton was determined about 750 to 800 °C (Fig. 7) (Tables 2a & b).

Geo-thermometry of Turkeh Dareh pluton based on zircon saturation
To determine the minimum temperature of zircon crystallization using saturation temperature [6, 7, 29, 30, 31, 32], specified conditions from Janousek & Saturnin [29] must exist in intrusion. The necessary conditions are that magma nature must be meta-aluminus, the absence of inherited zircons, zircon uniform distribution in entire rock and negative correlation between elements

Table 2. A) Morphological frequency distribution of zircon crystals on the basis of quantity and frequency of occurrence in Turkeh Dareh intrusion, B) zircon crystals emplacement within the area of 750 to 800°C on Pupin [2] diagram.
of zirconium and silica that these correlations are also true in the Turkeh Dareh pluton.

**Calculation of temperature based on Zircon saturation using Watson and Harrison Equation**

Watson and Harrison [6] confirmed the relationship between the solubility of zircon, temperature, and composition of the melt based on Equation 2:

\[
\ln D_{Zr} = \{12900/T (K) \} - 0.85 (M-1) - 3.80
\]  
(Equation 2)

In this equation, \( D_{Zr} \) is rate of zirconium concentration in zircon (500000 ppm) relative to the concentration of zirconium in molten (ppm). \( T \) represents the temperature in Kelvin and \( M \) is a cationic ratio depending on \( \text{SiO}_2 \) and aluminous melt calculated through Equation 3:

\[
M = [(Na + K + 2Ca)/(Al.Si)]
\]  
(Equation 3)

According to Equation provided after Watson and Harrison [6], zircon crystallization temperature of Turkeh Dareh intrusion was calculated 782 to 825 \( ^\circ \text{C} \) (Table 1).

Calculation of temperature based zircon saturation using Equation of Boehnke et al [7] revised the relationship between the solubility of zircon, temperature and composition of the molten provided from Watson and Harrison [6] and provided Equation 4 to estimate the saturation temperature of zircon (variables are like Equation 2):

\[
\ln D_{Zr} = \{(10108\pm 32)/T (K) \} - (1.16\pm 0.15) (M-1)
\]  
\[\pm (1.48 \pm 0.09)\]  
(Equation 4)

According to equation Boehnke et al [7], zircon crystallization temperature of Turkeh Dareh intrusion was calculated 702 to 800 \( ^\circ \text{C} \) (Table 1). According to the graph Zr versus \( \text{SiO}_2 \) of Watson and Harrison [6], the minimum crystallization temperature of Turkeh Dareh intrusion is 820 to 845 \( ^\circ \text{C} \) (Fig. 8).

To determine origin of Turkeh Dareh intrusion on the basis of the morphology of zircon, graph of Gehmich & Drost [18] was used. In this method, due to morphology of Turkeh Dareh intrusion zircons that are in the ranges of IIa, IIb, IIc, the magma origin of Turkeh Dareh intrusion is calc-alkaline and I-type (Fig. 9).

Observing distinct oscillatory zoning, typical of magmatic crystallization in Turkeh Dareh intrusion zircon indicate supersaturation of magma compared to zirconium from beginning of crystallization to final stages. The minimum zircon crystallization temperature in granitoides studies based on morphology is 700 to 850 \( ^\circ \text{C} \), based on zircon saturation temperature is 732 to 841 \( ^\circ \text{C} \) and based on whole rock geochemistry is 750 to 860 \( ^\circ \text{C} \). Considering the results of thermometer and supersaturation of granitic magma from beginning of crystallization relative to zirconium and as zircon is among the first crystalized minerals, so temperature range of 750 to 840 \( ^\circ \text{C} \) can be considered as minimum temperature of magma crystallization Turkeh Dareh intrusion. Growth of pyramids \( \{101\} \) of most zircons in the Turkeh Dareh intrusion represents origin of their calc-alkaline that it is consistent with the results of whole rock chemistry. Growth of prisms \( \{100\} \) in these zircons represents its high temperature index confirmed results of geothermometry. Based on mineralogy,
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geochemistry, and morphology of zircon Turkeh Dareh intrusion is I-type granite.

References

Figure 9. Pupin scheme showing the zircon morphology depending on (Na + K)/Al, and formation temperature [2] or on the content of U, Y, and P [7] with different trends: Legend: I = Al-rich anatetic granitoids; IIa,b,c = calc-alkaline granitoids; III = K-rich subalkaline granitoids; IV = alkaline granitoids [18].


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